

## PRECIPITATION

- is any dump or deposition in solid or liquid form that originated from the atmosphere.
- Most come from the sea.
- occurs when water vapor cools.
- When the air reaches saturation point (also know as condensation point and dew point) the water vapor condenses and forms tiny droplet of water.
- Precipitation form a $0.001 \%$ of the world's water resources


## Types of Precipitation

- HORIZONTAL PRECIPITATION
- Water vapour become frozen near the surface or directly on the surface
- Example : dew and fog
- VERTICAL PRECIPITATION
- Water vapour become freezing on the air and fall to the earth
- Example : rain and snow


## RAINFALL

- Rain is the major from precipitation in the form of water drops sizes greater than 0.5 mm.
- The maximum sizes of rain drop is about 6 mm based on the intensity, rainfall is classified as light rain ( $>2.5 \mathrm{~mm}-5 \mathrm{~mm} / \mathrm{hr}$ ), moderate rain ( $2.5-7 \mathrm{~mm} / \mathrm{hr}$ ) and heavy rain ( $>7.5 \mathrm{~mm} / \mathrm{hr}$ )



## Precipitation (Rain) Formation Process



## CYCLONE RAIN



## CYCLONE RAIN



## Hurricane \& Tropical Cyclone Formation



Fully formed tropical cyclone with an eye, eyewall, and circulating rain bands

## CONVECTIONAL RAIN



## CONVECTIONAL RAIN/ HUJAN PEROLAKAN

- Convectional precipitation results from the heating of the earth's surface that causes air to rise rapidly. As the air rises, it cools and moisture condenses into clouds and precipitation. This type of precipitation is common in the prairie provinces.


## Gonvectional

Precipitation

## Cooled Mir

## Condenses:

Stage 1 :
Stage 2
Stage 3 :
Stage 4 :

The sun heats the ground and warm air rises As the air rises it cools and water vapour condenses to form clouds
When the condensation point is reached large cumulonimbus clouds are formed Heavy rain storms occur. These usually include thunder and lightening due to the electrical charge created by unstable conditions.

## CONVECTIONAL RAIN/ HUJAN PEROLAKAN



## OROGRAPHIC PRECIPITATION



Prevailing wind

Moist, warm air rises Temperature $\sim 20^{\circ} \mathrm{C}$

Ocean

Stage $1 \quad$ : $\quad$ Warm wet air id forced to rise over high land
Stage 2 : As the air rises it cools and condenses. Clouds form and precipitation occurs.
Stage $3 \quad: \quad$ The drier air descends and warms
Stage 4 :
Any moisture in the air (e.g. cloud) evaporates

OROGRAPHIC PRECIPITATION results when warm moist air of the ocean is forced to rise by large mountains. As the air rises it cools, moisture in the air condenses and clouds and precipitation result on the windward side of the mountain while the leeward side receives very little

## RAIN MEASUREMENT


(a) Simple rain gauge

(b) Standard rain gauge

## (c) Tipping-bucket gauge




## RAIN MEASUREMENT



Modern Tipping Bucket Rain Gauge (Casella)


A commercial tipping bucket rain gauge
(Lambrech GmbH)



## Placement of Rain Gauges

Gauges are affected by wind pattern, eddies, trees and the gauge itself, therefore it is important to have the gauge located and positioned properly.

- 1 m above ground level is standard all gauges in a catchment should be the same height
- 2 to 4 times the distance away from an isolated object (such as a tree or building) or in a forest a clearing with the radius at least the tree height or place the gauge at canopy level


## Placement of Rain Gauges

- shielded to protect gauge in windy sites
- or if obstructions are numerous they will reduce the windspeed, turbulence and eddies.


Without Shield


With Shield

## Rain gauge with wind guards



## Placement of Rain Gauges

-For sloping ground the gauge should be placed with the opening parallel to the ground
-The rainfall catch volume $\left(\mathrm{mm}^{3}\right)$ is then divided by the opening area that the rain can enter


## Number and Distribution of Gauges

Need to consider:

- size of area
- prevailing storm type
- form of precipitation
- topography
- aspect
- season


## Number of Gauges

| Number of gauges | Size of area |
| :---: | :---: |
| 2 | 16 hectares |
| 3 | 40 hectares |
| 10 | $8 \mathrm{~km}^{2}$ |
| 15 | $16 \mathrm{~km}^{2}$ |
| 50 | $160 \mathrm{~km}^{2}$ |
| 300 | $1600 \mathrm{~km}^{2}$ |
| $1000+$ | $16,000 \mathrm{~km}^{2}$ |

- Snow is a form of precipitation within the Earth's atmosphere in the form of crystalline water ice, consisting of a multitude of snowflakes that fall from clouds. Since snow is composed of small ice particles, it is a granular material. It has an open and therefore soft structure, unless packed by external pressure. Snowflakes come in a variety of sizes and shapes. Types which fall in the form of a ball due to melting and refreezing, rather than a flake, are known as grapple, with ice pellets and snow grains as examples of graupel. Snowfall amount and its related liquid equivalent precipitation amount are determined using a variety of different rain gauges.

- Sleet refers to two distinct forms of precipitation:
- Rain and snow mixed, snow that partially melts as it falls. (UK, Ireland and Canada)
- Ice pellets, one of three forms of precipitation in a US-style "wintry mix", the other two being snow and freezing rain. (US)


## Sleet



Sleet forms when rain passes through a cold layer of air and freezes into ice pellets. This occurs most often in the winter when warm air is forced over a layer of cold air.


- Hail is a large frozen raindrop produced by intense thunderstorms, where snow and rain can coexist in the central updraft. As the snowflakes fall, liquid water freezzes onto them forming ice pellets that will continue to grow as more and more droplets are accumulated. Upon reaching the bottom of the cloud, some of the ice pellets are carried by the updraft back up to the top of the storm.
- Dew is water in the form of droplets that appears on thin, exposed objects in the morning or evening. As the exposed surface cools by radiating its heat, atmospheric moisture condenses at a rate greater than that at which it can evaporate, resulting in the formation of water droplets.
- When temperatures are low enough, dew takes the form of ice; this form is called frost (frost is, however, not frozen dew).
- Mist is a phenomenon of small droplets suspended in air. It can occur as part of natural weather or volcanic activity, and is common in cold air above warmer water, in exhaled air in the cold, and in a steam room of a sauna. It can also be created artificially with aerosol canisters if the humidity conditions are right.


## Rainfall Characteristic

- Depth
- Measurements in high-rainfall
- Measured in mm @ cm
- Measured by the hour, day, week, month and year Duration
- The period of time during which rain fell
- Rain time interval fall measured in minutes and hour
- Intensity
- Depth of rainfall per unit time
- Example: 30 mm in 3 hours; intensity $=10 \mathrm{~mm} / \mathrm{hr}$
- Rainfall Frequency
- The frequency of rainfall occurred
- Repeat for the rain
- This situation can be illustrated by the rainfall that is:
a) Uniform distribution: suitable for agriculture and reduce erosion
b) Distribution of single-mode: rainfall in the wet season
c) Distribution of dual mode


## Rainfall Intensity

- Intensity/heavy rain is the depth of rainfall per unit time
- The symbol is "i" and the units are mm/hr cm/hr @ in/hr
- Rainfall intensity equation :
i $=$ Depth rainfall
Time
Volume, $\mathrm{V}=$ Area $\times$ Depth

$$
=m^{3} @ k a^{3}
$$

## Rainfall Intensi Rainfall Intensity

|  | Intensity inches/hour (cm/hour) | Median diameter (millimeters) | Velocity of fall feet/second (meters/second) | Drops per second per square foot (square meter) |
| :---: | :---: | :---: | :---: | :---: |
| Fog | $\begin{gathered} 0.005 \\ (0.013) \end{gathered}$ | 0.01 | $\begin{gathered} 0.01 \\ (0.003) \end{gathered}$ | $\begin{gathered} 6,264,000 \\ (67,425,000) \end{gathered}$ |
| Mist | $\begin{gathered} .002 \\ (.005) \end{gathered}$ | . 1 | $\begin{gathered} .7 \\ (.21) \end{gathered}$ | $\begin{gathered} 2,510 \\ (27,000) \end{gathered}$ |
| Drizzle | $\begin{gathered} .01 \\ (.025) \end{gathered}$ | . 96 | $\begin{aligned} & 13.5 \\ & (4.1) \end{aligned}$ | 14(151) |
| Light rain | $\begin{gathered} .04 \\ (1.02) \end{gathered}$ | 1.24 | $\begin{aligned} & 15.7 \\ & (4.8) \end{aligned}$ | $\begin{gathered} 26 \\ (280) \end{gathered}$ |
| Moderate rain | $\begin{aligned} & .15 \\ & (.38) \end{aligned}$ | 1.60 | $\begin{gathered} 18.7 \\ (5.7) \end{gathered}$ | $\begin{gathered} 46 \\ (495) \end{gathered}$ |
| Heavy rain | $\begin{gathered} .60 \\ (1.52) \end{gathered}$ | 2.05 | $\begin{aligned} & 22.0 \\ & (76 .) \end{aligned}$ | $\begin{gathered} 46 \\ (495) \end{gathered}$ |
| Excessive rain | $\begin{gathered} 1.60 \\ (4.06) \end{gathered}$ | 2.40 | $\begin{aligned} & 24.0 \\ & (7.3) \end{aligned}$ | $\begin{gathered} 76 \\ (818) \end{gathered}$ |
| Cloudburst | $\begin{gathered} 4.00 \\ (10.2) \end{gathered}$ | 2.85 | $\begin{aligned} & 25.9 \\ & (7.9) \end{aligned}$ | $\begin{gathered} 113 \\ (1,220) \end{gathered}$ |

Source: Lull, H.W., 1959, Soil Compaction on Forest and Range Lands, U.S. Dept. of Agriculture, Forestry Service, Misc. Publication No. 768

Methods used to convert point rainfall values at different rain gauge stations into an average value over a catchment include

## Arithmetical Average Method

Thiessen Polygon Method

Isohyetal Method

Arithmetical Mean Method
Thiessen Polygon Method
Isohyetal Method etc

## Arithmetic Average Method

- This technique calculates area precipitation using the arithmetic mean of all the point or area measurements considered in the analysis.

$$
\bar{P}=\frac{P_{1}+P_{2}+P_{3}+\ldots \ldots \ldots \ldots+P_{n}}{n}=\frac{1}{n} \sum_{1}^{n} P_{i}
$$

$P=$ average precipitation over the catchment area (for a giventime period) $P_{1}, P_{2}, P_{3}, \ldots \ldots \ldots . . . ., P_{n}$ are the precipitations in a given time period at stations $1,2,3, \ldots . ., n$ respectively within . the catchment

Example:

## Thiessen Polygon Method

If there are $n$ stations with $r a \inf$ all values $P_{1}, P_{2}, P_{3}, \ldots . . P_{n}$ and $A_{1}, A_{2}, A_{3}, \ldots . A_{n}$ are the areas of the respective Thiessen polygons, the average rainf all over the catchment $P$ is computed as
$\bar{P}=\frac{P_{1} A_{1}+P_{2} A_{2}+P_{3} A_{3}+\ldots \ldots \ldots \ldots \ldots+P_{n} A_{n}}{A_{1}+A_{2}+A_{3}+\ldots \ldots \ldots \ldots \ldots+A_{n}}=\sum_{1}^{n} P_{i} \frac{A_{i}}{A}$
$\frac{A_{i}}{A}$ is called the weightage factor

## Thiessen Polygon Method



| Rainfall <br> Station | A <br> Thiessen Constant <br> (Fractional Area) | B <br> Rainfall in Each Polygon | A*B <br> Weighted average |
| :---: | :---: | :---: | :---: |
| A | 0.144 | 33.4 | 4.81 |
| B | 0.110 | 34.2 | 3.76 |
| C | 0.104 | 33.3 | 3.46 |
| D | 0.133 | 34.5 | 4.59 |
| E | 0.132 | 35.0 | 4.62 |
| F | 0.113 | 37.0 | 4.18 |
| G | 0.064 | 37.3 | 2.39 |
| H | 0.105 | 35.5 | 3.73 |
| I | 0.103 | 35.0 | 3.60 |
| Sum Total | $\mathbf{1 . 0 0}$ |  | $\mathbf{3 5 . 1 4}$ |



## Isohyetal Method

- Isohyet - It is a line joining points of equal rainfall magnitude
- The catchment area is drawn to scale and the rain gauge stations are marked on it
- The recorded rainfall values for which areal average is to determined are marked at the respective stations
- Neighbouring stations outside the catchment are also considered
- Taking point rainfall values as the guide, ishohyets of different rainfall values are drawn (similar to drawing contours based on spot levels)

The area between adjacent isohyets is measured using a planimeter.

- If isohyets go out of the catchment, the catchment boundary is used as the bounding line. It is assumed that the average value of rainfall indicated by two isohyets acts over the inter isohyetal area
- This method is considered superior to the previous methods when the number of raingauge stations are large


If the rainf all values corresponding to the isohyets are $P_{1}, P_{2}, P_{3}, \ldots . P_{n}$ and $A_{1}, A_{2}, A_{3}, \ldots \ldots A_{n-1}$ are the corresponding nt er isohyetal areas, the average rainf all over the catchment

P is computed as
$\overline{\mathrm{P}}=\frac{A_{1}\left(\frac{P_{1}+P_{2}}{2}\right)+A_{2}\left(\frac{P_{2}+P_{3}}{2}\right)+\ldots \ldots \ldots \ldots \ldots+A_{n-1}\left(\frac{P_{n-1}+P_{n}}{2}\right)}{A_{1}+A_{2}+A_{3}+\ldots \ldots \ldots \ldots \ldots+A_{n-1}}$
where, $A_{1}+A_{2}+A_{3}+\ldots \ldots \ldots \ldots \ldots+A_{n-1}=A$

| Isohyets <br> Upper | Lower | A <br> Mean Rainfall <br> On Area | Brea Between <br> Isohyets (ha) | A*B <br> Weighted Mean <br> Rainfall |
| :---: | :---: | :---: | :---: | :---: |
| 38 | 37 | 37.5 | 130 | 4875 |
| 37 | 36 | 36.5 | 150 | 5475 |
| 36 | 35 | 35.5 | 300 | 10650 |
| 35 | 34 | 34.5 | 450 | 15525 |
| 34 | 33 | 33.5 | 200 | 6700 |
| Total |  |  | $\mathbf{1 2 3 0}$ | $\mathbf{4 3 2 2 5}$ |

Mean Rainfall $=35.14$ i.e. from (43225/1230)


Catchment boundary

Normal ratio method (NRM) is used when the normal annual precipitation at any of the index station differs from that of the interpolation station by more than $10 \%$. In this method, the precipitation amounts at the index stations are weighted by the ratios of their normal annual precipitation data in a relationship of the form:

Where:

$$
P_{m}=\frac{1}{n} \sum_{i=1}^{n}\left(\frac{N_{m}}{N_{i}} P_{i}\right)
$$

$P \mathrm{~m}=$ precipitation at the missing location
$\mathrm{Pi}=$ precipitation at index station
$N m=$ average annual rain at 'missing data'
gauge
$\mathrm{Ni}=$ average annual rain at gauge
$N=$ number of rain gauges
is used to calculate quickly the average rainfall for hourly time intervals. The method takes into account the nonlinearity of the rainfall distribution over the basin area. The basin is divided in a network of square elements with a fixed step
size. The rainfall is calculated in each node using interpolation, as a function of the measured values at the adjacent stations.


## Procedure:

- Divide area around gauge of interest into four quadrants
- Using records at nearest station in each quadrant
- Compute missing precipitation amount:

$$
P_{m}=\frac{1}{\sum_{i=1}^{4} 1 / X_{i}}\left(\sum_{i=1}^{4} \frac{P_{i}}{X_{i}}\right)
$$

Where:
$\mathrm{Pi}=$ rainfall recorded by gauge $i$ $X i=$ distance from gauge $i$ to missing data point

## double-mass curve

Consistency of Precipitation Data A double-mass curve is used to check the consistency of a rain gauge record:

- compute cumulative rainfall amounts for suspect gauge and check gauges
- plot cumulative rainfall amounts against each other (divergence from a straight line indicates error)
- multiplying erroneous data after change by a correction factor $k$ where
$k=\frac{\text { gradient of line before change }}{\text { gradient of line after change }}$

- http://gs.riverdale.k12.or.us/~simonc17/panama/
- http://www.geography.hunter.cuny.edu/~tbw/wc.notes/5.cond.precip/sleet formation. htm
- http://ww2010.atmos.uiuc.edu/\(Gh\)/guides/mtr/cld/prcp/rnhl.rxm
- http://en.wikipedia.org/wiki/Dew
- http://www.egr.msu.edu/~northco2/BE481/HW1.htm
- http://theconstructor.org/water-resources/mean-precipitation-calculation-over-anarea/4500/
- http://www.aboutcivil.com/measurement-of-precipitation.html
- http://echo2.epfl.ch/VICAIRE/mod 1a/chapt 3/main.htm
- http://www.greenstone.org/greenstone3/nzdl;jsessionid=30A7616109C05E596B19423862F261F1 ?a=d\&c=cdl\&d=HASH20df2bee4882705a3d73c1.11.2.np\&sib=1\&p.s=ClassifierBrowse\&p.sa=\&p. $\mathrm{a}=\mathrm{b}$

